

**Lagrangian sampling of the Atlantic Water
pathways in the Nordic Seas:
Surface drifters and RAFOS floats**

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Three-dimensional Lagrangian view of the NwAc system

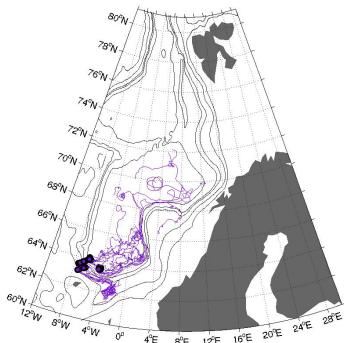
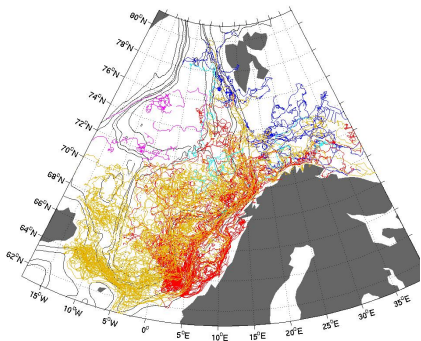
Surface drifters and rafos floats

- 306 surface drifters (15m, SVP), satellite positioning, 1990-2009
Global Drifter Programme + POLEWARD=118
- 22 intermediate-depth, acoustic positioning, isobaric RAFOS floats (ca. 800m),
2004-2005 (PATHMIX, Sjøiland et. al. 2008)
- 23 subsurface RAFOS floats (ca. 200m, IFF, Rossby et. al. 2009) - not yet
implemented

Analysis

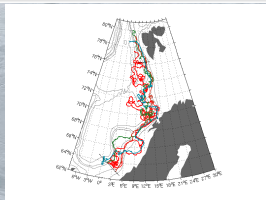
- Lagrangian signal at different depths - compare the flow statistics: dispersion,
travel times, speeds, topographic steering
- data-related problems and conceptual difficulties...

306 SVP (15m) + 22 RAFOS (800m)



3 challenging problems within Lagrangian analysis

- 1 limited data coverage
- 2 inherent flow inhomogeneities
- 3 non-stationarity

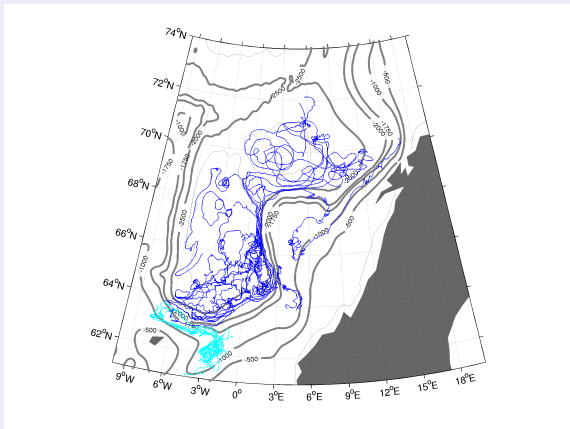


3 approaches within Lagrangian analysis of SVP/RAFOS

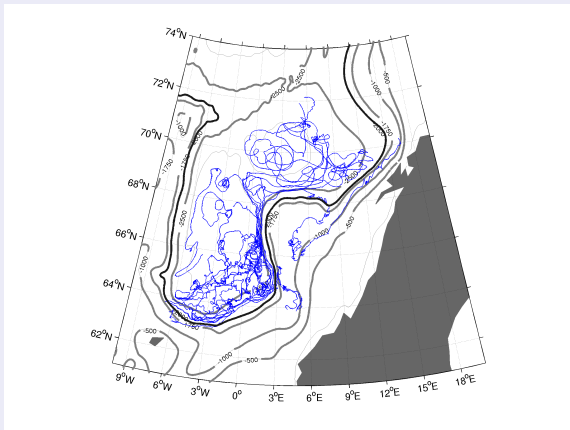
- 1 compare buoys passing through a location at any given time (nonstationarity, data paucity)
- 2 compare the statistics on buoy in certain regions (inhomogeneity)
- 3 bin (spatially-average) data over certain regions

The story begins in the Iceland Faroe Ridge....

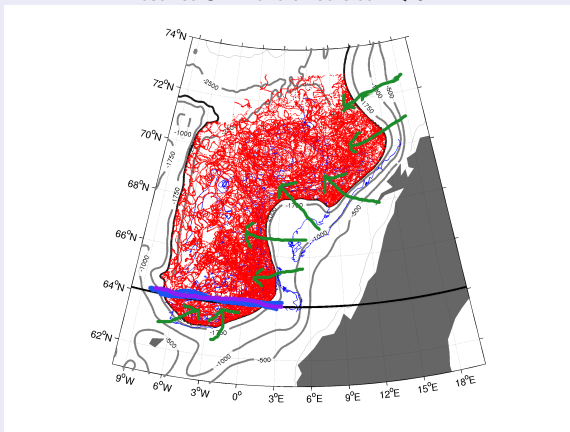
Søiland et. al. (2008) - fate determined by the topography:
Deep Group (13, deployed $< -1750\text{m}$) & Shallow Group (9, deployed $> -1750\text{m}$)



Deep Group extended: 13 Deep Group + 1 = 14.
Restrict area: NOR + LOF (2000m isobath)

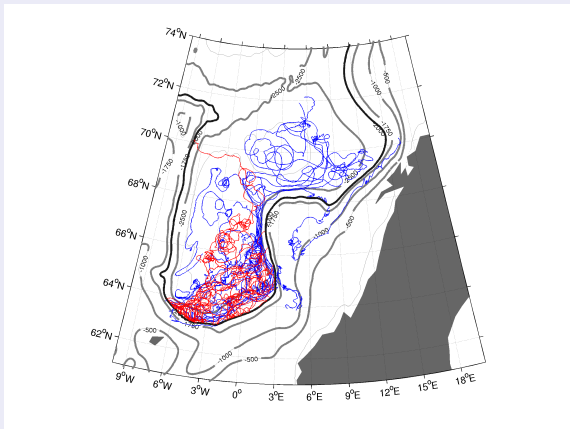


126 surface drifters - different enter site (NwAC-EB!) - not Lagrangian-comparable
Restrict SVP entrance site: $< 64^{\circ}\text{N}$

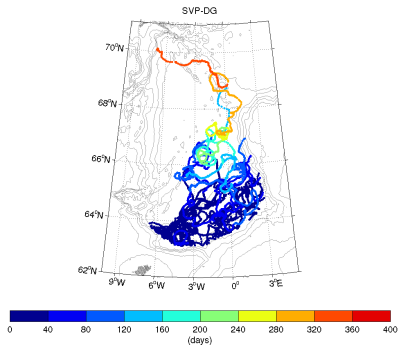
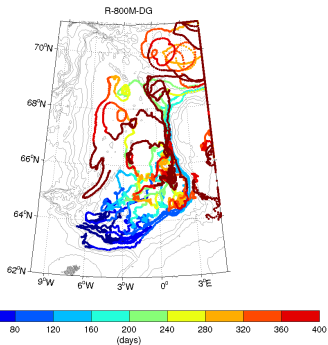


A study case: the Norwegian Basin

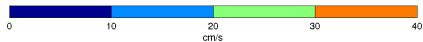
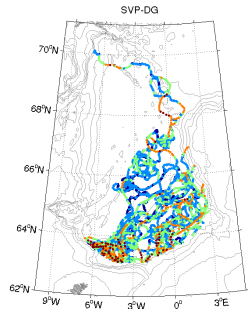
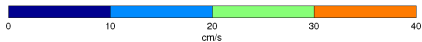
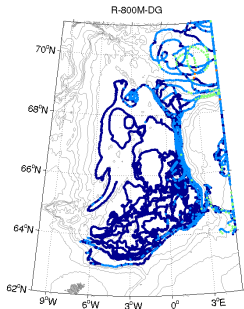
48 surface drifters & 14 rafos floats



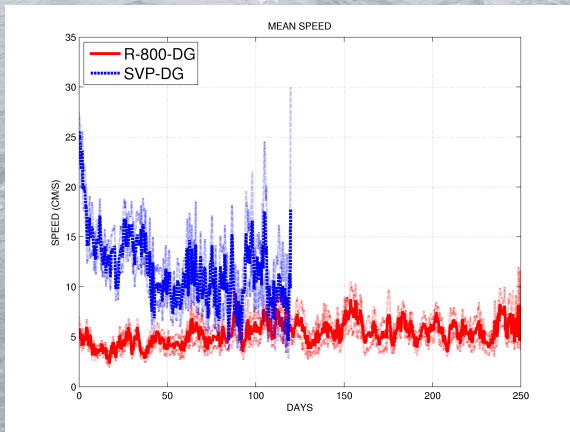
The Norwegian Basin: Pathways and time scales



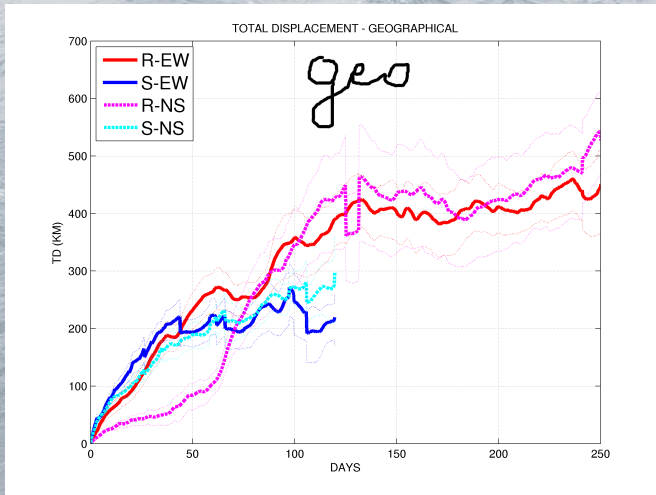
The Norwegian Basin: Speeds

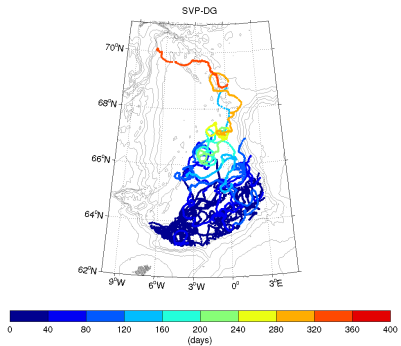
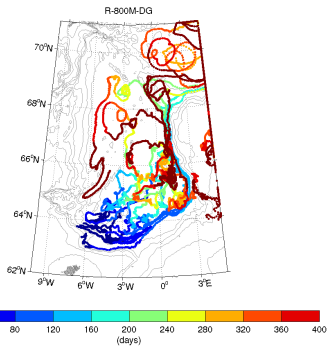
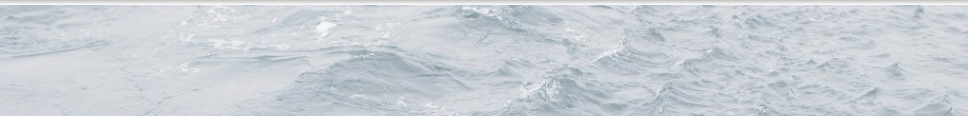


The Norwegian Basin: Speeds



The Norwegian Basin: Total displacement



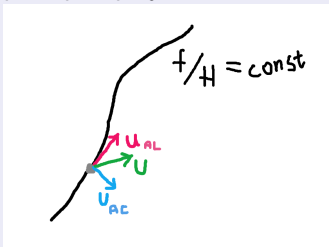


f/H analysis

- Nordic Seas - weakly stratified
- barotropic potential vorticity $PV \approx f/H$ flows tend to conserve it
- contribution of f one order smaller
- principle: project the velocities on f/H contours

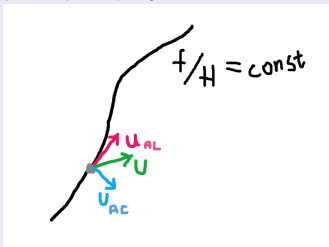
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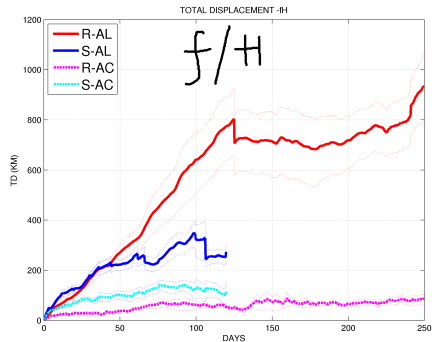
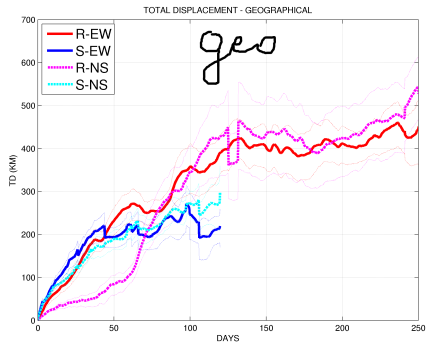
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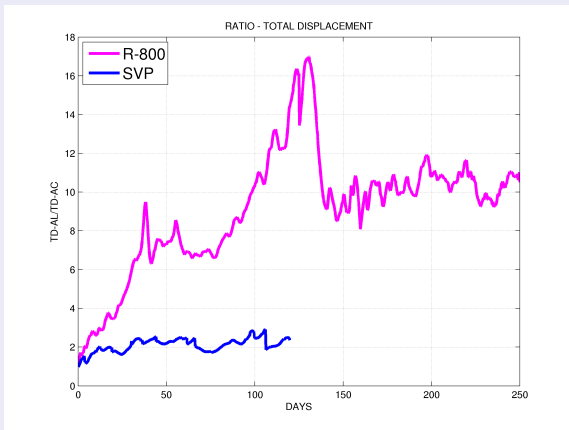


- integrate displacements
- scrutinize the anisotropy

The Norwegian Basin: Total displacement anisotropy



Degree of anisotropy - calculate the ratios



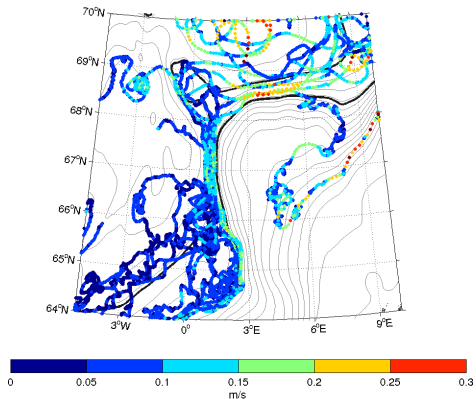
Unreliable statistics?

Idea: angle between the f/H -contours and drifter/float tracks

Vøring Margin

Rafos at VM (11)

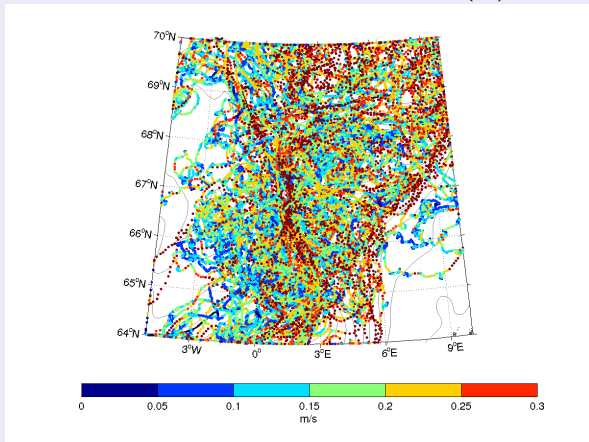
Rafos: strong topographic steering at the Voring Margin (66-68 N, 0.2.5 E)
Søiland et. al. (2008)



Vøring Margin

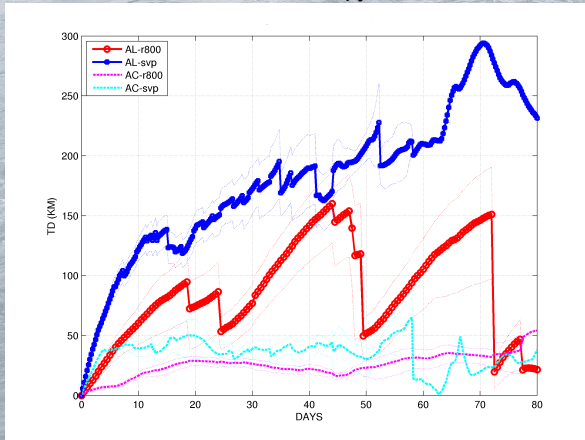
Surface drifters at VM

Steering av SVP: mentioned by Orvik and Niiler (2002)
All surface drifters recorded at the VM (58)



Vøring Margin

fH-anisotropy:



Ratio: 3.8 (rafos), 2.6 (surface drifters)

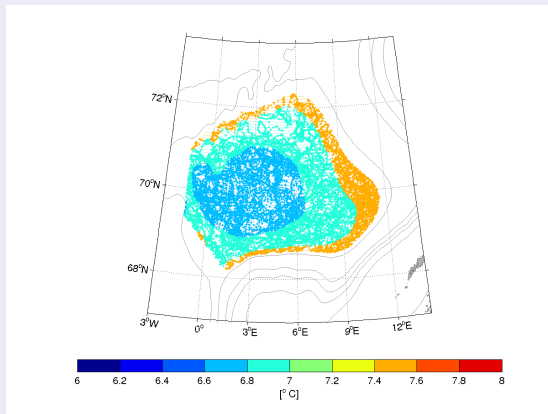
Vøring Margin

Transport at VM

- Rafos, 500-2500m: $z=2000\text{m}$, $u \approx 0.1 \text{ m/s} \times 30\text{km} \rightarrow 6 \text{ sv}$ (Sjøiland et. al., 2008)
- more complicated for surface drifters (spread around the VM)
- mean speed ($\sqrt{u^2 + v^2} \approx 0.25 \text{ m/s}$) and mean along-fH speed (u_L)
- u_L constitutes ca. 90% $|u|$
- between 2200 and 2900 isobath (ca. 18km):
($uv=3.6 + 1.4=5\text{sv}/uL=3.6 + 0.9=4.5 \text{ sv}$)
- between 2000 and 3000 isobath (ca. 35km):
($uv=6.3 + 2.3 = 8.6/uL=6.3 + 1.75=8.1 \text{ sv}$)
- depends on smoothness of topography/distribution of data - needs more careful calculations

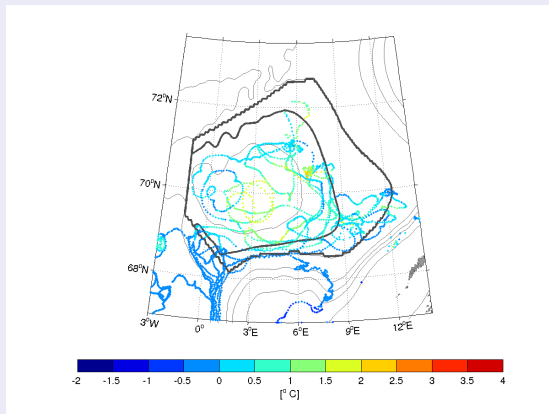
Lofoten Basin - temperature signal

Surface signal - binned in f/H contours



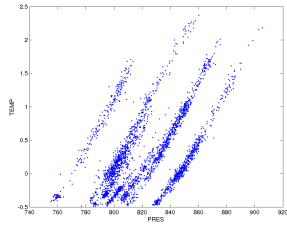
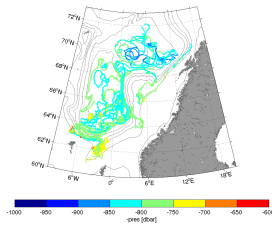
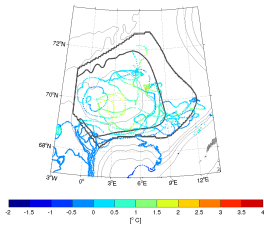
Lofoten Basin - temperature signal

Rafos signal at 800m



Lofoten Basin - temperature signal

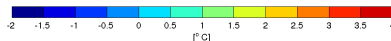
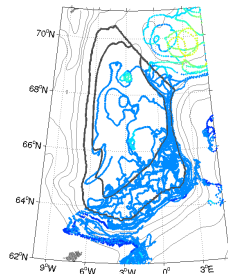
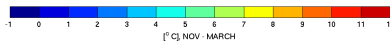
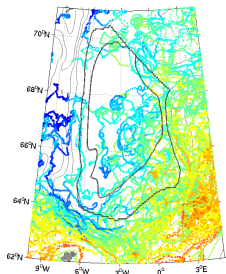
Rafos signal at 800m



AW-layer deepening (Orvik, 2004) + MLD deepening in winter (Nilsen and Falck, 2006)
The role of eddies in cooling (lateral mixing)
Vertical mixing (deep eddies/MLD - Nilsen and Falck, 2006)

The Norwegian Basin

Binning of temperature in fH-contours does not work in the Norwegian Basin
the Arctic Front!



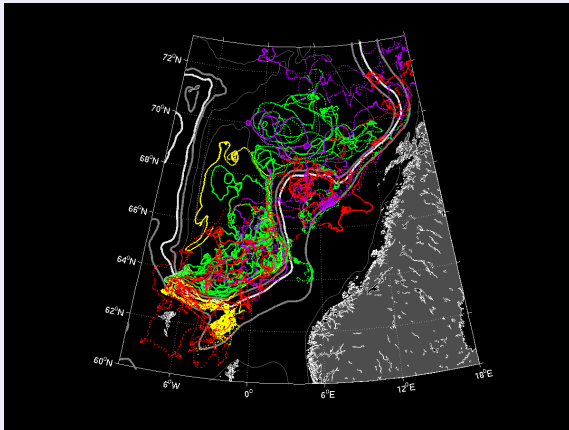
The role of eddies!

Initial conclusions

- SVP + RAFOS data sets offer a powerful possibility to obtain a 3d-Lagrangian view of the circulation in the NS
- First attempt to quantify the evident topographic steering in the NwAC WB-EB-NCC system has been made (displacement anisotropy, binning of temperature signal in f/H -contours)
- The topographic steering in the SE Norwegian Basin appears stronger at depths than at the surface
- Strong topographic control at the VM at all depths (barotropic flow) - high transports?
- Further evidence on AW deepening in the Lofoten Basin, horizontal eddy mixing and vertical mixing (to be explored in more detail..)
- a reliable statistics to assess the differences is still to be found (work in progress)
- try to integrate the RAFOS data from 200m depth into the analysis

The story in the Iceland Faroe Ridge revisited

RAFOS: DG (13, deployed $< -1750\text{m}$) & SG (9, deployed $> -1750\text{m}$)
SVP "deployed" $dr = (+/- 0.02^\circ)$: "DG" (12) & SG (4)



Idea: increase dr , check destinations.